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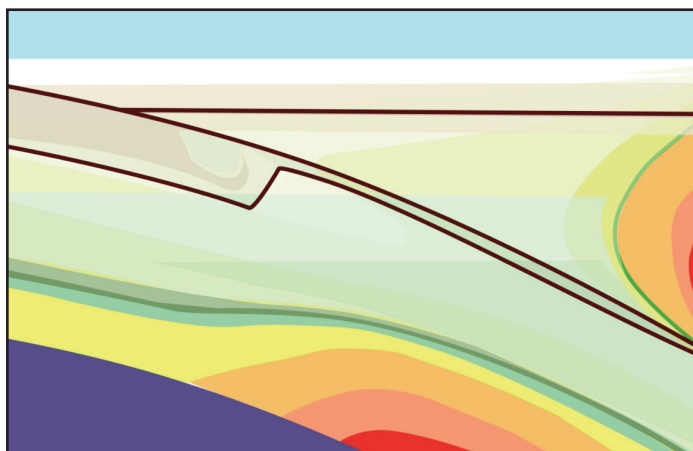
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Article abstract

In northwest Donegal, Ireland, a large number of coeval appinitic (hornblende-plagioclase-rich) plutons and lamprophyre dykes occur around the Ardara pluton, a granitic satellite body and one of the oldest phases of the ca. 428–400 Ma composite Donegal Batholith. The appinite units form a bimodal (mafic–felsic) suite in which hornblende is the dominant mafic mineral and typically occurs as large prismatic phenocrysts within a finer grained matrix. Lamprophyre dykes are mafic in composition with a geochemistry that is very similar to that of the mafic appinite bodies. Both mafic rocks are subalkalic, with calc-alkalic and tholeiitic tendencies, and show trace element abundances indicating that the mantle source was contaminated by subduction zone fluids. $^{40}\text{Ar}/^{39}\text{Ar}$ analysis of hornblende separated from two samples of appinite yield mid-Silurian (434.2 ± 2.1 Ma and 433.7 ± 5.5 Ma) cooling ages that are interpreted to closely date the time of intrusion. Hence, according to the available age data, the appinite bodies slightly predate, or were coeval with, the earliest phases of the Donegal Batholith. Sm–Nd isotopic analyses yield a range of initial ϵNd values (+3.1 to –4.8 at $t = 435$ Ma) that, together with trace element data, indicate that the appinitic magmas were likely derived from melting of metasomatized sub-continental lithospheric mantle and/or underplated mafic crust, with only limited crustal contamination during magma ascent. The appinitic intrusions are interpreted to have been emplaced along deep-seated crustal fractures that allowed for mafic and felsic magma to mingle. The magmas are thought to be the products of collisional asthenospheric upwelling associated with the closure of Iapetus and the ensuing Caledonian orogeny, either as a result of an orogen-wide delamination event or as a consequence of more localized slab break-off.

ANDREW HYNES SERIES: TECTONIC PROCESSES



Age, Geochemistry and Origin of the Ardara Appinite Plutons, Northwest Donegal, Ireland

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SUMMARY

In northwest Donegal, Ireland, a large number of coeval appinitic (hornblende-plagioclase-rich) plutons and lamprophyre dykes occur around the Ardara pluton, a granitic satellite

body and one of the oldest phases of the ca. 428–400 Ma composite Donegal Batholith. The appinite units form a bimodal (mafic–felsic) suite in which hornblende is the dominant mafic mineral and typically occurs as large prismatic phenocrysts within a finer grained matrix. Lamprophyre dykes are mafic in composition with a geochemistry that is very similar to that of the mafic appinite bodies. Both mafic rocks are sub-alkalic, with calc-alkalic and tholeiitic tendencies, and show trace element abundances indicating that the mantle source was contaminated by subduction zone fluids. ⁴⁰Ar/³⁹Ar analysis of hornblende separated from two samples of appinite yield mid-Silurian (434.2 ± 2.1 Ma and 433.7 ± 5.5 Ma) cooling ages that are interpreted to closely date the time of intrusion. Hence, according to the available age data, the appinite bodies slightly predate, or were coeval with, the earliest phases of the Donegal Batholith. Sm–Nd isotopic analyses yield a range of initial εNd values (+3.1 to –4.8 at *t* = 435 Ma) that, together with trace element data, indicate that the appinitic magmas were likely derived from melting of metasomatized sub-continental lithospheric mantle and/or underplated mafic crust, with only limited crustal contamination during magma ascent. The appinitic intrusions are interpreted to have been emplaced along deep-seated crustal fractures that allowed for mafic and felsic magma to mingle. The magmas are thought to be the products of collisional asthenospheric upwelling associated with the closure of Iapetus and the ensuing Caledonian orogeny, either as a result of an orogen-wide delamination event or as a consequence of more localized slab break-off.

RÉSUMÉ

Dans le nord-ouest du Donegal, en Irlande, un grand nombre de plutons appinitiques (riches en hornblendes ou en plagioclases) et de dykes de lamprophyres contemporains se retrouvent autour du pluton d'Ardara, un corps satellite granitique et l'une des phases les plus anciennes du batholite composite de Donegal, âgé d'environ 428–400 Ma. Les unités de l'appinite forment une suite bimodale (mafique–felsique) dans laquelle la hornblende est le minéral mafique dominant et se présente généralement sous forme de grands phénocristaux prismatiques au sein d'une matrice à grains plus fins. Les dykes de lamprophyres ont une composition mafique dont la géochimie est très similaire à celle des corps d'appinite mafique. Les deux roches mafiques sont subalcaliques, avec des tendances calco-alcalines et tholéiitiques, et elles montrent des teneurs en élé-

ments traces indiquant que la source du manteau a été contaminée par des fluides de zone de subduction. L'analyse $^{40}\text{Ar}/^{39}\text{Ar}$ des hornblendes provenant de deux échantillons d'appinite donne des âges de refroidissement du Silurien moyen ($434,2 \pm 2,1$ Ma et $433,7 \pm 5,5$ Ma) qui sont interprétés comme étant proches de la date de l'intrusion. Par conséquent, selon les données d'âge disponibles, les corps d'appinite sont légèrement antérieurs ou contemporains des toutes premières phases du batholite de Donegal. Les analyses isotopiques Sm–Nd aboutissent à une gamme de valeurs ϵNd initiales (+3,1 à -4,8 à $t = 435$ Ma) qui, associées aux données des éléments traces, indiquent que les magmas appinitiques sont probablement dérivés de la fusion d'un manteau lithosphérique sous-continental métasomaté et / ou d'une croûte mafique sous-plaquée, avec une contamination crustale limitée lors de l'ascension du magma. Les intrusions appinitiques sont interprétées comme s'étant mises en place le long de fractures profondes de la croûte qui ont permis au magma mafique et au magma felsique de se mélanger. On pense que les magmas sont les produits de la remontée (upwelling) asthénosphérique collisionnelle associée à la fermeture de l'océan Iapetus et à l'orogène calédonienne qui s'ensuit, soit à la suite d'un délaminage à l'échelle de l'orogène, soit à la suite d'une rupture plus localisée de la plaque.

Traduit par la Traductrice

INTRODUCTION

In many parts of the Caledonian orogenic belt of Scotland and Ireland, major granitoid bodies, such as the composite Donegal Batholith in NW Ireland, are associated with a distinct suite of relatively small, hornblende-rich plutons formed from water-rich magmas that are 'peri-batholithic' in that they occur around the periphery of the composite batholith (Pitcher and Berger 1972). Known collectively as the appinite suite after the district of Appin in Scotland (e.g. Walker 1927; Bowes and McArthur 1976; Hamidullah and Bowes 1987), these intrusive rocks, which range from mafic to felsic in composition, are unusual in their mineralogy and texture. Those of mafic composition are medium- to coarse-grained melanocratic rocks composed of large, idiomorphic hornblende crystals in a matrix of feldspar and minor quartz (e.g. Pitcher and Berger 1972; Wright and Bowes 1979). Those of intermediate to felsic composition are also rich in idiomorphic hornblende and include syenite, monzonite, diorite and granodiorite. The suite is also characterized by the widespread occurrence of associated breccia intrusions and brecciated metasedimentary rocks.

The appinite suite in both Scotland and Donegal is commonly associated with lamprophyre dykes that show a similar peri-batholithic relationship with respect to major granitoid batholiths (Atherton and Ghani 2002). This temporal and spatial association has been widely interpreted to reflect a genetic linkage between the appinite suite, the lamprophyre dykes and the granitoid batholiths (e.g. Pitcher and Berger 1972), all of which were collectively emplaced towards the end of the Scandian orogeny (e.g. Siegesmund and Becker 2000; Atherton and Ghani 2002).

In Ireland, appinite bodies comprise a sub-group of Caledonian igneous intrusions best exposed around the Ardara pluton at the southwestern end of the composite Donegal Batholith (Fig. 1). The Ardara pluton is one of several intrusive bodies emplaced along the periphery of the Main Donegal Granite (Pitcher and Berger 1972). Available geochronological data suggest that the composite Donegal Batholith contains discrete phases ranging in age from 428 to 400 Ma and that the Ardara pluton is among the oldest of these phases (Condon et al. 2004).

In the Scottish Caledonides, rocks of the appinite suite are also spatially and temporally associated with 430–408 Ma granitoid magmatism (e.g. Neilson et al. 2009). These appinite plutons are geochemically distinct from coeval mafic and granitoid rocks of other associations (Fowler and Henney 1996; Atherton and Ghani 2002). They are typically calc-alkaline and have compositional affinities with shoshonite (basaltic, K-rich trachyandesite composed of olivine, augite and plagioclase phenocrysts), features suggesting emplacement in a volcanic arc tectonic setting (Thompson and Fowler 1986; Macdonald et al. 1986; Fowler 1988). Their distinctive geochemistry and the timing of their emplacement (broadly synchronous with arc–continent collision) has been interpreted to reflect melting of a previously metasomatized lithospheric mantle as a consequence of slab break-off following the termination of subduction (e.g. Atherton and Ghani 2002; Ghani and Atherton 2008; Neilson et al. 2009; Cooper et al. 2013). In addition, recent proposals that large batholiths preferentially form during the waning stages of collision and consequent slab failure (Hildebrand and Whalen 2014a, b; Hildebrand et al. 2018) indicate that rocks of the appinite suite have the potential of contributing significantly to our understanding of collisional processes and the voluminous granitoid magmas with which they are associated.

Despite their potential importance to understanding the genesis and tectonic setting of the Caledonian intrusive rocks, published geochemical data for the appinitic rocks of Donegal are sparse and include only a limited suite of petrogenetically indicative trace elements (e.g. French 1966; Hall 1967), and the precise age of their intrusion is unknown. To further our understanding of these enigmatic rocks and their association with voluminous granitoid batholiths, this paper aims to clarify the intrusive age and petrogenesis of this suite through geochronological ($^{40}\text{Ar}/^{39}\text{Ar}$ on hornblende), geochemical and isotopic (Sm–Nd) analysis of the appinite bodies associated with the Ardara pluton.

GEOLOGICAL SETTING

The Donegal Batholith of NW Ireland (Fig. 1), made famous by the classical work of Pitcher and Berger (1972), is made up of a suite of Caledonian granitic plutons emplaced into Laurentian crust in the mid-Silurian–early Devonian (428–400 Ma, U/Pb: Condon et al. 2004; 418–397 Ma, Rb/Sr: Halliday et al. 1980; O'Connor et al. 1982, 1987), during the final closure of the Iapetus Ocean (Ghani and Atherton 2008). During the development of the Caledonian orogen, the Laurentian margin

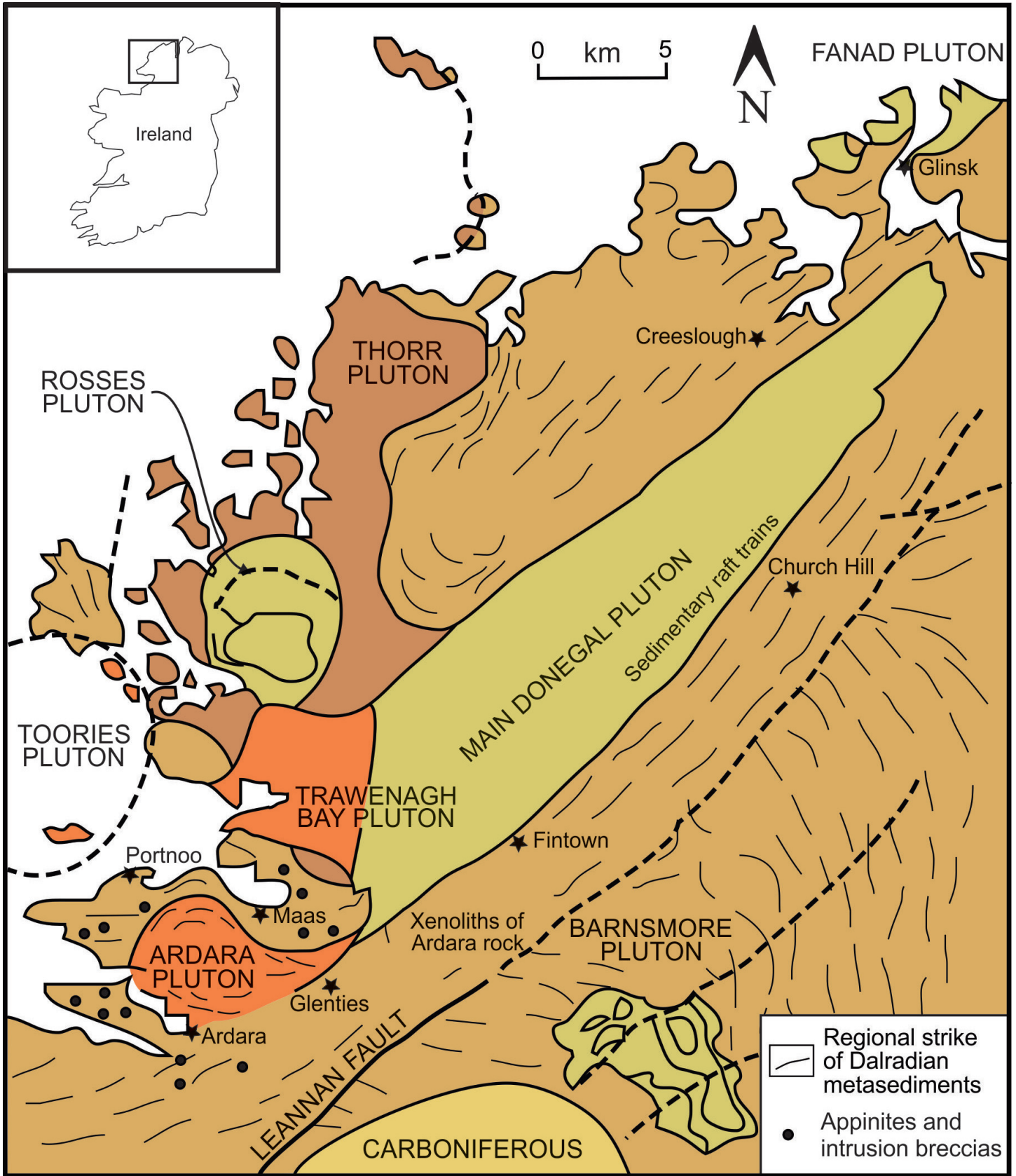


Figure 1. Simplified geological map of the composite Donegal Batholith in northwest Ireland (from Ghani and Atherton 2008).

was affected by two distinct orogenic episodes (e.g. van Staal et al. 1998). The first was the short-lived, mid-Ordovician Grampian event (Dewey and Shackleton 1984; Oliver 2001; Chew and Strachan 2014) linked to arc formation and accretion associated with the subduction of Iapetus oceanic lithosphere below the Laurentian margin at ca. 470–460 Ma. The second and main episode records the final closure of Iapetus and the collision of the Laurentia margin with Baltica-Ganderia (Scandian orogeny) and Avalonia (Acadian orogeny) in mid-Silurian to early Devonian time (e.g. Condon et al. 2006; Chew and Strachan 2014). In Ireland, the resulting Iapetus suture strikes NE–SW from Clogherhead in County Louth to the Dingle Peninsula in County Kerry. Ireland's crust northwest of this suture was originally derived from Laurentia, whereas the crust to the southeast comprises terranes that were originally derived from Gondwana (Todd et al. 1991).

The Donegal Batholith is a composite, predominantly granitoid body made up of the Thorr, Fanad, Trawenagh Bay, Rosses, Barnesmore, Toories and Ardara plutons in addition to the Main Donegal Batholith (Pitcher and Berger 1972; Vernon and Patterson 1993) (Fig. 1). These plutons were emplaced into the Dalradian Supergroup country rock at mid-crustal levels (Siegesmund and Becker 2000) and display distinct differences regarding their shape, internal structures, contact aureoles and emplacement mechanisms. The Dalradian Supergroup itself was deposited along the eastern margin of Laurentia during the late Neoproterozoic and early Cambrian and comprises a thick sequence of lithologically diverse metasedimentary and mafic volcanic rocks (Ghani and Atherton 2008). Dalradian lithologies in Donegal include mica schist, dolomitic schist, and striped and banded siliceous quartzite (Ghani and Atherton 2008).

The Main Donegal Granite (MDG) is the largest pluton within the Donegal Batholith, forming an elongate, steep-walled body with sharp margins that cross-cut intensely deformed Dalradian polymetamorphic schist, quartzite and metabasite (Pitcher and Read 1960; Hutton 1982; Stevenson et al. 2008). Although predominantly granitic, the MDG is internally sheared and varies from mafic to felsic in composition. It is characterized by a foliation and contains inclusions of Dalradian country rock oriented parallel to the length of the pluton (Berger 1971). These features are reportedly the result of stresses superimposed upon the pluton after its emplacement but before complete cooling and consolidation had occurred (French 1976; Hutton and Alsop 1996). The high-K, calc-alkaline granite and granodiorite that make up the MDG are thought to have been emplaced during a phase of major NE–SW sinistral strike-slip faulting that characterizes the final phase of Iapetus closure in the British Caledonides (Atherton and Ghani 2002). The MDG is temporally and spatially associated with both lamprophyre dykes and appinitic bodies, the latter being most common adjacent to the Ardara pluton at its southwestern margin (Fig. 1).

Ardara Pluton

Marginal to the MDG, but also contributing to xenoliths within it (Price 1997), the smaller Ardara pluton is a concentrically

zoned granodiorite–quartz monzodiorite body (Fig. 2), the outer zone of which is quite strongly deformed (French 1966; Vernon and Patterson 1993). It is a complex, roughly circular body about 10 km in diameter and exhibits sharp contact zones (French 1966). During emplacement, the metasedimentary Dalradian host rocks underwent contact metamorphism and ductile deformation, the latter thought to have provided most of the space for the pluton (Siegesmund and Becker 2000). There were three phases of intrusion; an early outer zone of quartz diorite, an inner zone of coarse potassium-feldspar megacrystic granodiorite and a later core composed of equigranular granodiorite. Unpublished U/Pb data for the Ardara pluton date its intrusion ca. 427 Ma (Condon et al. 2004).

The Ardara pluton has been regarded as a classic example of emplacement by expansion (ballooning) due to the injection of significant volumes of magma into the centre of the body after initial intrusion (Pitcher and Berger 1972). The bulk of the foliation within the pluton is magmatic (apart from a thin rind of solid-state deformation), suggesting that the expansion was accompanied by magmatic flow (Harmon et al. 1984). Flattened microgranitoid enclaves, concentric magmatic flow foliation, and compositional zoning within the pluton are consistent with expansion of magma prior to crystallization (Vernon and Patterson 1993). Mafic enclaves with variable modes of occurrence and texture resemble the lithologies in the appinitic intrusions. Some enclaves within the pluton were sufficiently rigid to fragment without any field evidence of mixing or mingling (Atherton and Ghani 2002). These field relationships suggest the intrusion of appinitic magmas preceded the Ardara pluton, although there may be some temporal overlap between early stages of pluton emplacement and the lamprophyre dykes (Atherton and Ghani 2002).

Appinitic Suite

The Ardara pluton is surrounded by a large number of small appinitic intrusions and associated lamprophyre dykes, which are also thought to have been intruded during the final stages of the Caledonian orogeny (Atherton and Ghani 2002). The rocks of appinitic intrusions constitute the appinitic suite and are most commonly found in small stocks, sheets, bosses and dykes, commonly closely associated with late to post-tectonic granitoid intrusions (e.g. Fowler and Henney 1996). The suite primarily occurs along the southern edge of the MDG and within 3 km of the Ardara pluton, where it additionally forms small irregular pipes or vents, clustered in large numbers around the plutons and as mafic enclaves within them.

Rocks of the appinitic suite are typically coarse-grained and porphyritic with green to brown hornblende in a groundmass with approximately equal proportions of plagioclase and potassium-feldspar (Hall 1967). The appinitic intrusions are predominantly lenticular masses, 100–550 m in width, and most have irregular margins (French 1966) with steeply inclined contacts that truncate regional Dalradian structures and related foliations. Several appinitic masses north of the pluton exhibit a foliation defined by the alignment of felsic minerals interpreted to have formed during their emplacement

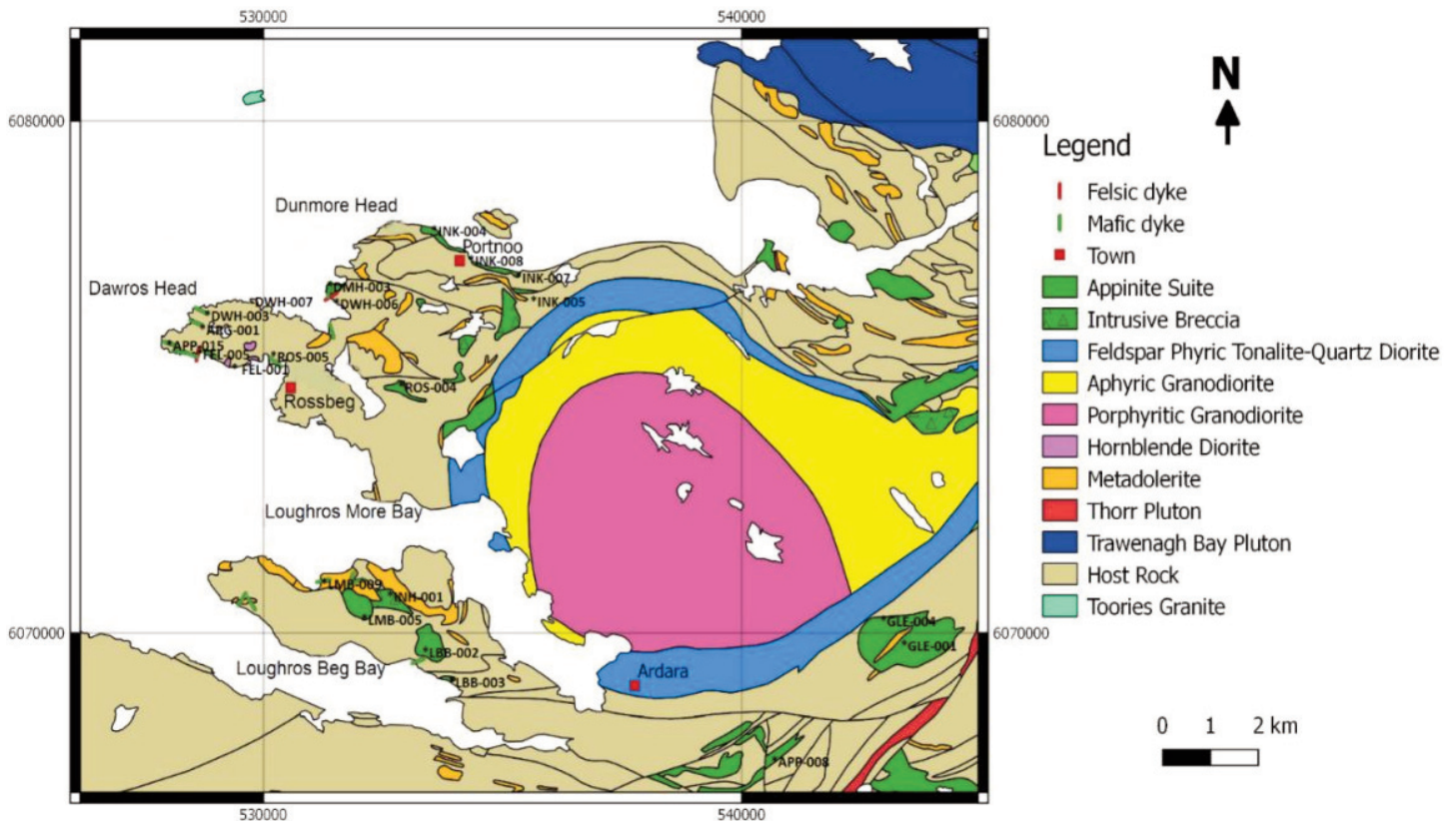


Figure 2. Geological map of the Ardara pluton showing sample locations (after Geological Survey of Ireland 2014). See Table 1 for details of locations.

and cooling (Bowes and McArthur 1976). The presence of xenoliths of appinitic rocks and lamprophyre within the MDG suggests that both were intruded prior to the emplacement of this pluton (French 1976). However, while the Ardara pluton has an early marginal phase lithologically identical to one of the rocks of the appinite suite, and both the appinitic intrusions and lamprophyre dykes crop out close to the pluton, neither cut the pluton itself (French 1976).

Many appinitic bodies are accompanied, particularly along their margins and in roof zones, by masses of disrupted metasedimentary rocks interpreted as intrusion breccias (French 1966). These breccia masses occur in near-vertical conduits that are elliptical in cross section with long axes up to 12 m. The clasts in the breccia are inferred to have been transported upwards to their present positions and comprise fragments of carbonate and quartzitic rocks that are typical of the Dalradian host rock (French 1966).

The appinite suite is made up of coeval plutonic and hypabyssal rocks that range from mafic (gabbro) to intermediate-felsic (diorite, granodiorite) in composition and contain prismatic hornblende in a groundmass of plagioclase, with or without quartz (Fig. 3a, b). The hornblende crystals are typically 3–5 mm in length and, in some cases, are in grain boundary contact with each other.

Although mingling of felsic and mafic lithologies within the appinite suite is rare, some local evidence of hybridization has been observed (Hall 1967). The relationship between the different lithologies is complex with evidence for multiple

intrusion and *in situ* differentiation, as well as complex interactions with the surrounding country rock (Platten 1991). Pitcher and Berger (1972) observed that some outcrops are dominated by intrusion breccia and explosion breccia characterized by angular fragments of quartzite and calc-silicate rock within an appinitic matrix. These features, together with the dominance of hornblende and the wide range of textures, suggest that crystallization of the appinitic magmas occurred under high water vapour pressure (Rock 1991).

Lamprophyre Dykes

The rocks of the appinite suite, along with the breccia intrusions associated with them, intrude and are intruded by lamprophyre dykes (Elsdon and Todd 1989). Lamprophyre (Fig. 3c) occurs as dykes and sills that intruded the Dalradian host rocks and is characterized by phenocrysts of hornblende and phlogopite with lesser amounts of clinopyroxene. The groundmass is diverse and includes plagioclase, biotite, pyroxene and amphibole. Lamprophyre variants include hornblendite, consisting mostly of idiomorphic hornblende within a relatively fine-grained matrix dominated by plagioclase. Less commonly, lamprophyre also contains interstitial quartz (Pitcher and Berger 1972).

⁴⁰Ar/³⁹Ar THERMOCHRONOLOGY

Analytical Methods

In order to date the emplacement of the Ardara appinite suite, ⁴⁰Ar/³⁹Ar thermochronology was performed at the

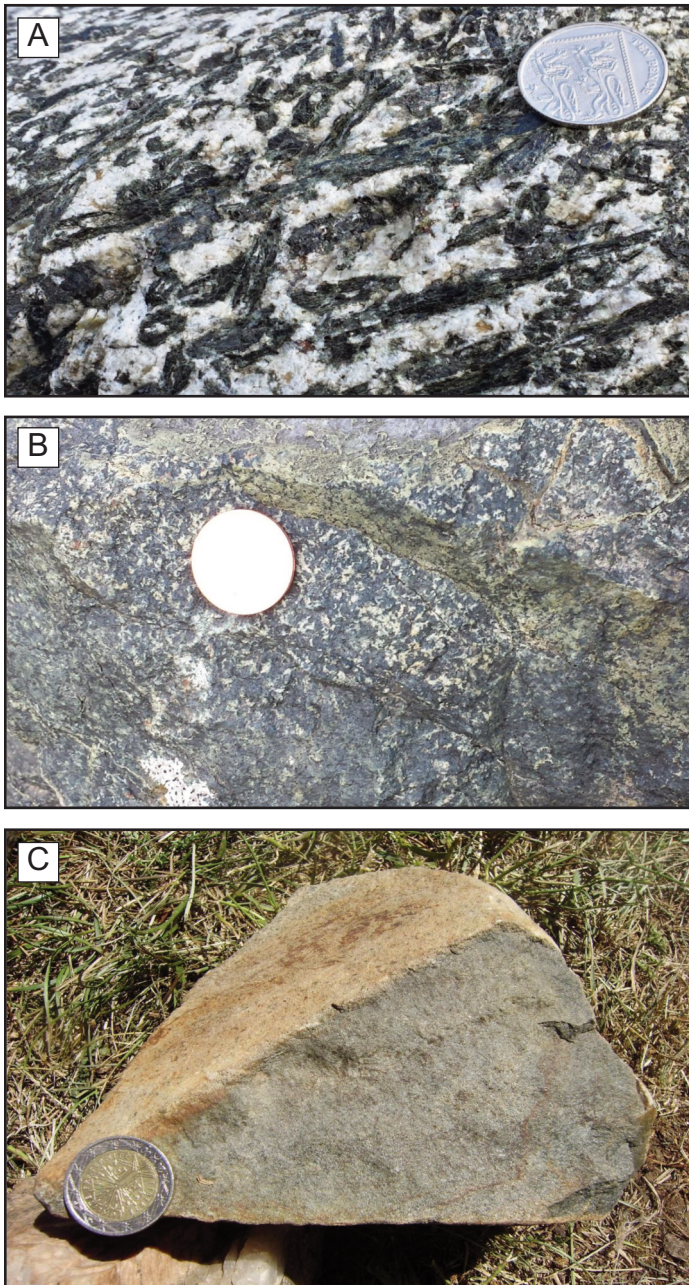


Figure 3. Field photographs of (a) mafic rock of the appinite suite with prismatic amphibole grains within a plagioclase matrix, (b) intermediate-felsic rock of the appinite suite with amphibole clumps within a plagioclase matrix, and (c) a lamprophyre dyke with fine-grained biotite and hornblende grains within a plagioclase-rich matrix.

geochronology laboratory at Queen's University in Kingston, Ontario, on amphibole separated from two representative samples (INK-004 and GLE-004). Electron microprobe analyses and X-ray images of representative hornblende from each sample are shown in the supplemental files SF-1 to SF-5. The amphibole analyses show variations from 39.7 to 42.7 wt.% in SiO_2 , 11.8 to 13.9 wt.% in Al_2O_3 , 10.8 to 12.4 wt.% in CaO , 2.1 to 4.03 wt.% in TiO_2 and 0.59 to 1.24 wt.% K_2O . Mean Ca/K ratios are between 11.1 and 13.3. Because the water and halogen contents of the amphiboles are unknown, the amphibole

formulae are calculated to 23(O), and the $\text{Fe}^{2+}/\text{Fe}^{3+}$ ratio determined after the method described by Pe-Piper (1988). According to the classification of Leake et al. (2004), the amphiboles are calcic as they all have $(\text{Ca}+\text{Na})^{\text{B}} \geq 1.00$ and $\text{NaB} > 0.50$ apfu (atoms per formula unit). The amphiboles typically have Si^{IV} between 5.9 and 6.4, with Al^{IV} from 2.1 to 1.6 and $\text{Mg}/(\text{Mg}+\text{Fe}^{2+})$ between 0.6 and 0.9. The analyses indicate that the amphibole compositions are typical of tschermakitic hornblende of igneous origin (Leake 1978; Leake et al. 2004).

The sample locations are shown on Figure 2 and listed in Table 1. Both samples were hand-crushed and single grains of hornblende cleavage fragments free of visible overgrowths were selected for analysis. Mineral separates and flux-monitors (standards) were wrapped in Al foil. The resulting disks were stacked vertically into an 8.5 cm long and 2.0 cm diameter Al irradiation capsule, and then irradiated with fast neutrons in position 8C of the McMaster Nuclear Reactor (Hamilton, Ontario) for a duration of 86.4 h (at 2.5 MWh). Packets of flux monitors were located at ~ 0.5 cm intervals along the irradiation container and the J-value for each of the samples was determined by least-squares, second-order polynomial interpolation using weighted means of replicate analyses of splits of the bracketing monitors position in the capsule.

The samples were loaded into flat-bottomed pits in a copper sample-holder and placed beneath the ZnS view-port of a small, bakeable, stainless-steel chamber connected to an ultra-high vacuum purification system. Following bake-out at 105°C , a 30 W New Wave Research MIR 10-30 CO_2 laser with a faceted lens was used to heat the samples for 3 minutes at increasing percent power settings (2% to 45%; beam diameter 3 mm). After purification using hot and cold SAES C50 getters (for 5 minutes), the evolved gas was admitted to a MAP 216 mass spectrometer, with a Baur Signer source and an analogue electron multiplier (set to a gain of 100 over the Faraday detector).

Measured Ar isotope peak heights were extrapolated to zero-time and corrected for discrimination using an $^{40}\text{Ar}/^{36}\text{Ar}$ atmospheric ratio of 298.56 (Lee et al. 2006) and measured ratios of atmospheric Ar. Blanks, measured routinely, were subtracted from the subsequent sample gas fractions. The extraction blanks were typically $< 10 \times 10^{-13}$, $< 0.5 \times 10^{-13}$, $< 0.5 \times 10^{-13}$, and $< 0.5 \times 10^{-13} \text{ cm}^{-3}$ STP for masses 40, 39, 37, and 36, respectively. ^{39}Ar and ^{37}Ar were corrected for radioactive decay during and after irradiation. Corrections were made for neutron-induced ^{40}Ar from potassium, ^{39}Ar and ^{36}Ar from calcium, and ^{36}Ar from chlorine (Roddick 1983; Onstott et al. 1991). Dates and errors were calculated using the procedure of Dalrymple et al. (1981) and the ^{40}K decay constant of Min et al. (2000; $5.463 \times 10^{-10} \text{ 1/y}$). Plateau and inverse isotope correlation dates were calculated using ISOPLOT v. 3.60 (Ludwig 2008). A plateau is herein defined as 3 or more contiguous steps containing $> 50\%$ of the ^{39}Ar released, with a probability of fit > 0.01 and $\text{MSWD} < 2$. If the contiguous steps contain $< 50\%$ of the ^{39}Ar released, it is referred to as a plateau segment.

Quoted errors represent the analytical precision at 2σ , assuming that the error in the age of the flux monitor is zero.

Table 1. Lithology, location and analysis of samples selected for geochemical analysis from the Ardara appinite plutons and coeval lamprophyre dykes, Donegal, Ireland.

<i>Sample ID</i>	<i>Rock Type</i>	<i>Location</i>	<i>Northing</i>	<i>Easting</i>	<i>Analysis</i>
Appinite Suite					
<i>APP-008</i>	Mafic Dyke	Ardara	54.7538	-8.3819	REE
<i>APP-015</i>	Mafic Dyke	Dawros Head	54.8333	-8.4552	REE
<i>LBB-003</i>	Mafic Dyke	Loughros Beg Bay	54.7698	-8.4816	REE, Sm–Nd
<i>INK-004</i>	Mafic Dyke	Portnoo	54.8475	-8.4801	REE, Sm–Nd, Ar–Ar
<i>AGR-001</i>	Mafic Dyke	Dawros Head	54.7337	-8.4332	REE, Sm–Nd
<i>GLE-001</i>	Mafic Dyke	Glenties	54.8223	-8.277	REE, Sm–Nd
<i>GLE-004</i>	Mafic Dyke	Glenties	54.7477	-8.2946	REE, Sm–Nd, Ar–Ar
<i>ROS-004</i>	Mafic Dyke	Rosbeg	54.8226	-8.5296	REE
<i>INK-007</i>	Felsic Dyke	Portnoo	54.82575	-8.5603	REE
<i>INK-008</i>	Felsic Dyke	Portnoo	54.84329	-8.4717	REE, Sm–Nd
<i>FEL-005</i>	Felsic Dyke	Dawros Head	54.8254	-8.5544	REE, Sm–Nd
<i>DWH-007</i>	Felsic Dyke	Dawros Head	54.8331	-8.5502	REE
<i>INK-005</i>	Felsic Dyke	Portnoo	54.84436	-8.4724	REE
<i>DWH-006</i>	Intermediate	Dawros Head	54.83297	-8.5532	REE
<i>INH-001</i>	Intermediate Dyke	Loughros Beg Bay	54.82922	-8.5108	REE
Lamprophyre Dykes					
<i>FEL-001</i>	Mafic Dyke	Dawros Head	54.8254	-8.5505	REE, Sm–Nd
<i>LMB-009</i>	Mafic Dyke	Loughros More Bay	54.8365	-8.514	REE, Sm–Nd
<i>LBB-002</i>	Mafic Dyke	Loughros Beg Bay	54.77122	-8.4833	REE, Sm–Nd
<i>DMH-003</i>	Mafic Dyke	Dumore Head	54.8365	-8.5155	REE, Sm–Nd
<i>DWH-003</i>	Mafic Dyke	Dawros Head	54.8308	-8.5548	REE, Sm–Nd
<i>ROS-005</i>	Mafic Dyke	Rosbeg	54.82322	-8.529	REE
<i>LMB-005</i>	Mafic Dyke	Loughros More Bay	54.7870	-8.5143	REE

This precision is suitable for comparing within-spectrum variation and determining which steps form a plateau (e.g. McDougall and Harrison 1988, p. 89). The dates and J-values were referenced to GA1550 biotite (98.5 Ma; Spell and McDougall 2003; McDougall and Wellman 2011; recalculated relative to an FC sanidine age of 28.201 Ma; Kuiper et al. 2008).

Results

A summary of the $^{40}\text{Ar}/^{39}\text{Ar}$ data obtained is shown in supplemental file SF-6. Step heating of hornblende separates from these two samples resulted in release patterns with initially high ages, decreasing over the first few percent of the ^{39}Ar released, then increasing irregularly to robust well-defined plateaus. In this study, the ISOPLOT plateau dates represent the best estimate of the age of the samples. For both spectra, the plateau steps incorporate those with typical yield Ca/K ratios (12 to 15) calculated from the $^{37}\text{Ar}/^{39}\text{Ar}$ ratios comparable to those obtained from the electron microprobe analyses (11 to 13).

Sample INK-004 is from an appinitic intrusion into Dalradian host rock taken near the Portnoo coast. This sample contains abundant amphibole, biotite clusters and interpenetrating apatite in a matrix dominated by plagioclase and amphibole. It yielded a ten-step plateau with 90.5% of the ^{39}Ar released and a plateau date of 434.8 ± 1.2 Ma (Fig. 4a), and within error a concordant inverse isochron date of 434.2 ± 2.1 Ma (Fig. 4b). The inverse isochron included 12 of the 14 steps, 94.3% of the

^{39}Ar released, had a slightly non-atmospheric initial $^{40}\text{Ar}/^{36}\text{Ar}$ ratio (333 ± 18) and an MSWD = 0.63. The ages are essentially the same at each step except for those produced by the argon released (5%–10%) during the initial heating. During the first few steps, traces of loosely bound excess Ar caused the older dates in this part of the spectrum.

Sample GLE-004 is from an appinitic intrusion at Glenties in the contact aureole of the Ardara pluton ~200 m from the exposed contact. This sample is dominated by phenocrysts of olivine, amphibole and augite in a groundmass composed of phlogopite, biotite, plagioclase and orthoclase. The sample yielded a seven-step high-temperature plateau with a date of 440.1 ± 1.4 Ma (65.8% of the ^{39}Ar released) (Fig. 5a). Because the inverse isochron ratios for the high-temperature steps were tightly clustered, a reliable ISOPLOT Model 1 solution could not be obtained. A Model 2 solution using all points yielded a date of 433.7 ± 5.5 Ma (Fig. 5b) and an initial $^{40}\text{Ar}/^{36}\text{Ar}$ ratio of 405 ± 14 . The initial ratio is well above that for atmospheric argon and suggests that excess Ar was incorporated in the hornblende. As in Sample INK-004, the ages are essentially the same at each step except for those produced by the argon released (5%–30%) during the initial heating. During the first few heating steps, likely traces of excess Ar caused the ages to appear about 50–80 million years older.

Interpretation

The Ar closure temperature for hornblende is variable and



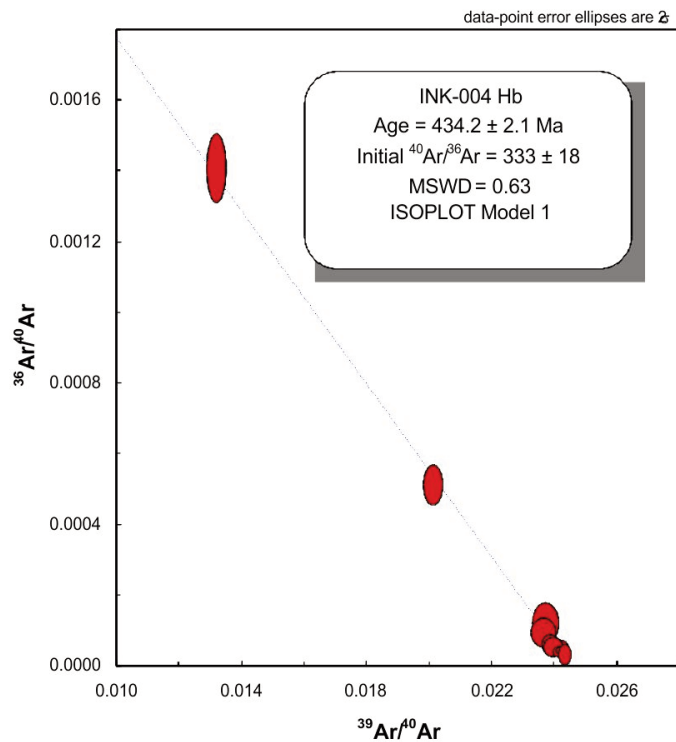
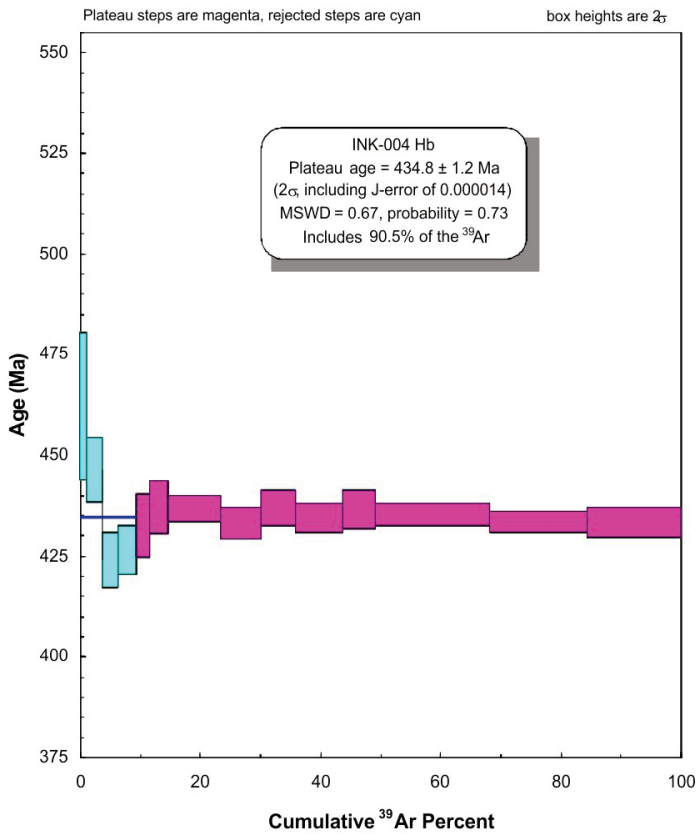


Figure 4. (a) $^{40}\text{Ar}/^{39}\text{Ar}$ age spectrum and (b) $^{40}\text{Ar}/^{39}\text{Ar}$ inverse isochron plot for sample INK-004.

depends on several compositional factors. Because the dated amphibole samples have Mg/(Mg+Fe) ratios in the range of the hornblende investigated by Harrison (1982; 0.6–0.8), it is

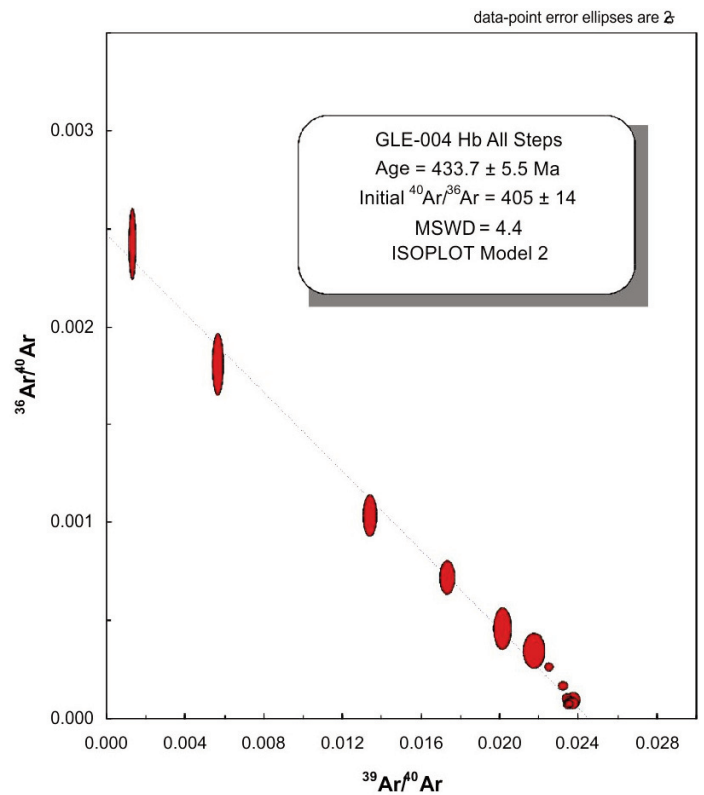
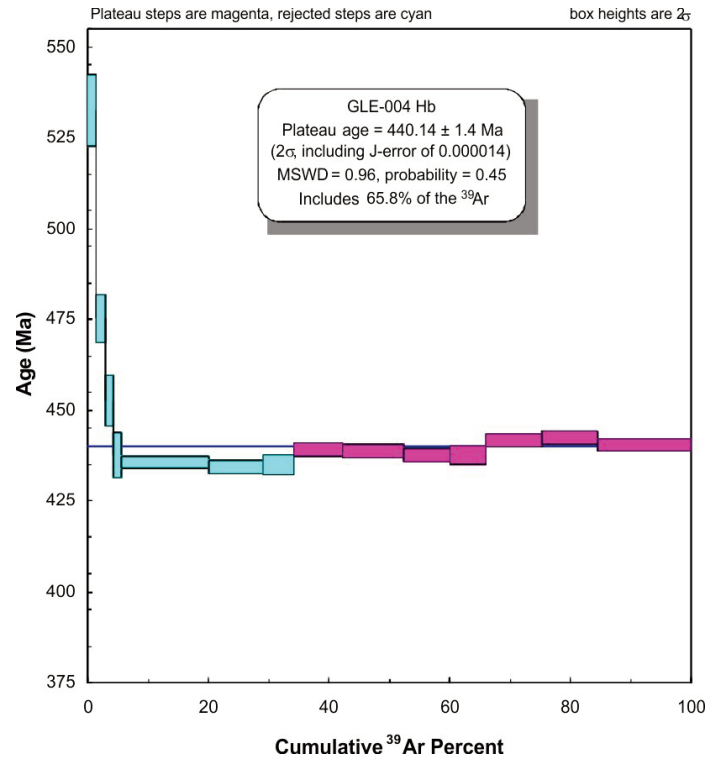


Figure 5. (a) $^{40}\text{Ar}/^{39}\text{Ar}$ age spectrum and (b) $^{40}\text{Ar}/^{39}\text{Ar}$ inverse isochron plot for sample GLE-004.

appropriate to calculate a closure temperature based on his experimental diffusion parameters. The relatively small size of both intrusions together with their mid-crustal level of

emplacement, suggests that cooling through the Ar closing temperature for hornblende was rapid. Using Dodson's (1973) iterative expression for closure temperature and a conservative cooling rate of 100°C/m.y. a probable minimum closure temperature of 545°C is obtained. Sample location INK-004 is ~ 2 km from the contact of the Ardara pluton and is not likely to have experienced any reheating from the Ardara pluton. GLE-004 is within ~ 500 m of the southern contact of the Ardara and may have been reheated. However, based on geothermometry and geobarometry in pelitic rocks, Kerrick (1987) modelled the thermal gradient across the aureole. Based on his contact temperature profile and pressures of 2–3 kbar, it is unlikely that GLE-004 would have been re-heated to a temperature > 500°C. The calculated closure temperature is significantly higher than the metamorphic temperatures experienced by the Dalradian metasedimentary rocks that host the Ardara appinite suite, which are at regional greenschist facies. The cooling age of 434.2 ± 2.1 Ma (2σ) obtained from the $^{40}\text{Ar}/^{39}\text{Ar}$ inverse isochron is consequently interpreted to have been acquired during rapid cooling of the appinitic magma. Given the relatively small size and mid-crustal levels of emplacement, both of which would suggest rapid cooling through the Ar closing temperature for hornblende, the ages are considered to date igneous emplacement closely. These data therefore indicate a Silurian (late Llandovery–early Wenlock) crystallization age for the Ardara appinite suite and, when combined with field relationships, suggest that emplacement of the suite slightly preceded emplacement of the earliest Donegal granite plutons, including the Ardara pluton (Fowler and Henney 1996; Atherton and Ghani 2002; Condon et al. 2004).

GEOCHEMISTRY

Twenty-two samples representative of the appinite suite and lamprophyre dykes were collected from all the larger appinitic bodies and several lamprophyre dykes surrounding the Ardara pluton. These were sent to the Bureau Veritas Commodities Canada, Ltd., Vancouver, British Columbia, for major and trace element analysis by X-ray fluorescence (XRF) using a Philips PW2400 X-ray spectrometer. The samples were obtained from fine- to coarse-grained appinitic intrusions and lamprophyre dykes intruding Dalradian host rock at Dunmore Head, Portnoo, Rossbeg, Loughros More Bay, Loughros Beg Bay and Ardara (Fig. 2). Sample locations are given in Table 1 and shown on Figure 2. Rare earth and selected trace elements were analyzed by Inductively Coupled Plasma Mass Spectrometry (ICP-MS), also at the Bureau Veritas Commodities Canada, Ltd., using a Fisons/Applied Research Laboratories 8420+ wavelength dispersive X-ray spectrometer. A subset of the powdered samples was additionally analyzed for Sm–Nd isotopic values at Memorial University using thermal ionization mass spectrometry.

Whole-rock Analysis

Samples of the appinite suite were taken from mafic enclaves within the Ardara pluton and from dykes up to 5 m in width that cut the Dalradian host rocks. Selected major and trace ele-

ment data for these samples are given in supplementary file SF-7. Loss-on-ignition (LOI) values are variable, ranging from 1.9 to 2.9 wt.% in mafic rocks and 1.1 to 2.6 wt.% in intermediate–felsic rocks of the appinite suite, and from 49.5 to 57.2 wt.% in the lamprophyre dykes.

Appinite Suite

Mafic rocks of the appinite suite have SiO_2 content ranging from 45.8–53.3 wt.%, MgO values from 5–16 wt.%, magnesium numbers (Mg#), $(100 \times \text{MgO} / (\text{MgO} + 0.9\text{FeO}_{\text{tot}}))$ that range from 22 to 57 and variable Al_2O_3 contents (13.0–22.0 wt.%). Intermediate–felsic members of the suite have higher SiO_2 concentrations (64.8–72.1 wt.%) but a similar range in Al_2O_3 (13.3–19.1 wt.%) and much lower MgO values (0.28–0.78 wt.%).

Petrographic evidence for alteration in rocks of the appinite suites necessitates the emphasis on concentrations of less mobile high-field strength elements (HFSE) and rare earth elements (REE) instead of the more mobile large ion lithophile elements (LILE) in the interpretation of petrogenesis and tectonic setting (e.g. Winchester and Floyd 1977). The mobility of alkalis (Na_2O , K_2O) is evident from their scatter on Harker variation diagrams (Fig. 6). The Zr/Ti vs Nb/Y diagram (Pearce 1996) is used to replace the total alkali–silica diagram because Zr/Ti and Nb/Y are respective proxies for the silica and alkali contents (Fig. 7). The mafic rocks of the suite are sub-alkalic and plot in the basaltic andesite to basalt fields. On $\text{FeO}_{\text{tot}}/\text{MgO}$ vs FeO and $\text{FeO}_{\text{tot}}/\text{MgO}$ vs SiO_2 (Miyashiro 1974) discrimination diagrams (Fig. 8a,b), the mafic rocks show a limited range in SiO_2 over a large range in $\text{FeO}_{\text{tot}}/\text{MgO}$, a feature that is typical of a tholeiitic differentiation trend, but the low $\text{FeO}_{\text{tot}}/\text{MgO}$ ratio exhibited by several mafic samples is typical of calc-alkaline suites.

The mafic rocks are also characterized by generally high MgO (4.1 to 16.7 wt.%), Cr_2O_3 (0.01 to 0.06 wt.%), Ni (98–288 ppm), Ba (233–528 ppm), and V (179–462 ppm) abundances as well as elevated LILE- and light rare-earth elements (LREE), and slightly depleted heavy rare-earth elements (HREE). The high concentrations of Mg, Ni, Cr and Ba are indicative of a mantle source (Atherton and Ghani 2002). On chondrite-normalized REE diagrams, the mafic rocks display a moderate negative Eu anomaly indicating fractionation of plagioclase (Fig. 9a). On multi-element plots (Pearce 1983), the rocks also display moderately negative anomalies in HFS elements such as Nb, P, Zr and Ti (Fig. 10a,b). The overall slope, a general source enrichment of LILE (Fig. 9b), and the negative anomalies in HFS elements are characteristic of a mantle source in a subduction zone setting. The mafic rocks additionally have generally high Zr/Y, a characteristic typical of within-plate tholeiitic basalt. In the Ti/1000 vs V diagram (Fig. 11), the mafic rocks plot in the fields of mid-ocean ridge (MORB) and back-arc basin (BAB) basalt.

On Th/Yb and Ce/Yb vs Ta/Yb plots (Fig. 12a,b), the vertical axis detects subduction components so that rocks formed in a coeval arc environment, or those contaminated by continental crust plot above typical mantle values. The rocks of the appinite suite display calc-alkalic to shoshonitic compositions



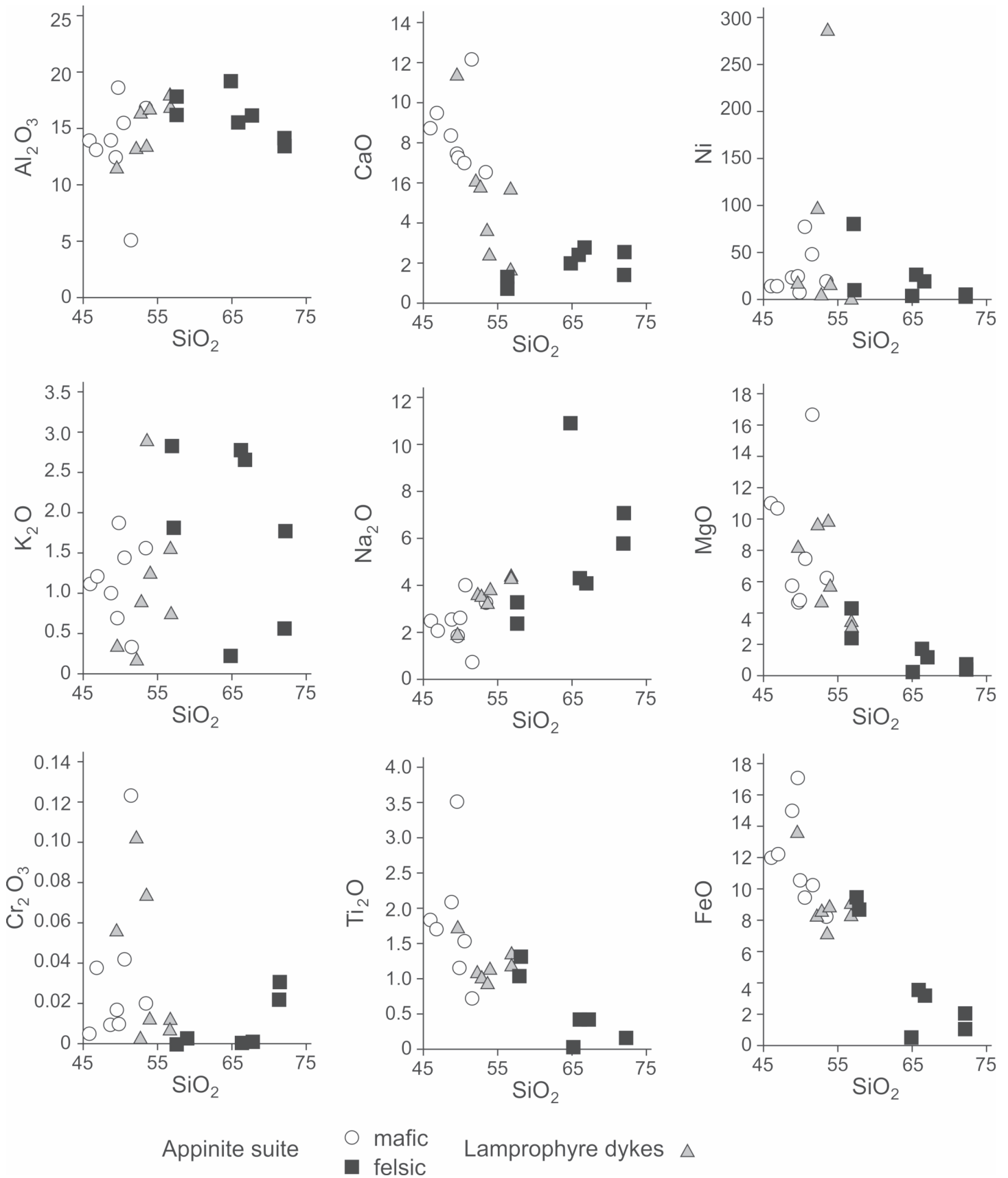


Figure 6. Harker diagrams for major element oxides for the Donegal appinitic suite.

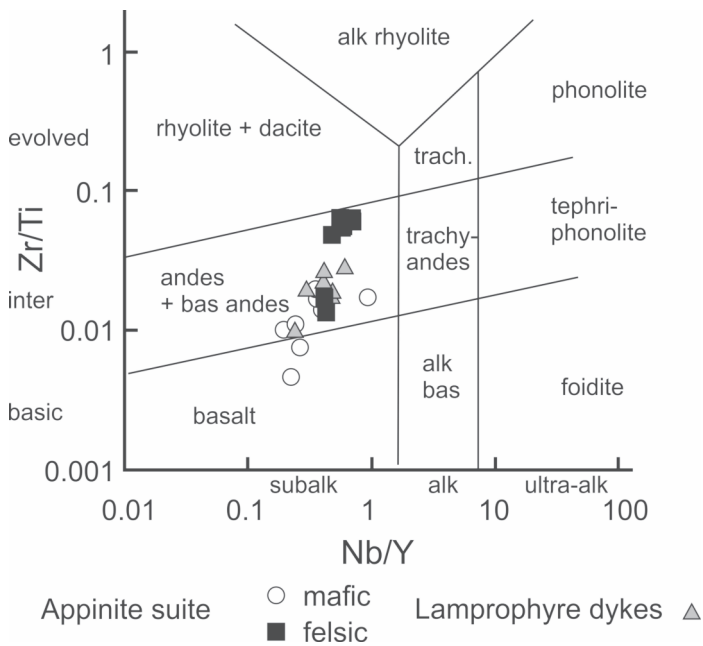


Figure 7. Zr/Ti vs Nb/Y diagram for classifying the Donegal appinites (Pearce 1996 after Winchester and Floyd 1977).

on these plots. On the La/Yb vs Th/Yb plot (Fig. 13), the mafic and intermediate–felsic rocks of the appinite suite display island arc to continental margin arc characteristics. As arc magmas are enriched in Th and depleted in Ta, they plot nearer the Hf–Th join compared to MORB or within-plate basalts on the Hf–Th–Ta diagram (Fig. 14). The rocks of the appinite suite plot as calc-alkaline basalt and have higher Th than typical island arc tholeiite (IAT), plotting closer to the Th apex of this discrimination diagram.

Lamprophyre Dykes

The lamprophyre dykes are mafic to intermediate and range from 49.5 to 56.7 wt.% in SiO₂, 11.5 to 18.0 wt.% in Al₂O₃, and 1.9 to 4.4 wt.% in Na₂O. These rocks are also high in TiO₂, MnO, P₂O₅ and Cr₂O₃.

Using the FeO_{tot}/MgO vs SiO₂ and FeO_{tot}/MgO vs FeO (Miyashiro 1974; Arculus 2003) diagrams, the lamprophyre dykes display small ranges in SiO₂ and FeO contents over a large range in FeO_{tot}/MgO (Fig. 8a,b). The slope evident in this figure is typical of a tholeiitic differentiation trend.

The lamprophyre dykes are moderately enriched in the LREE and LILE relative to HREE. They display a moderate negative Eu anomaly indicating fractionation of plagioclase (Fig. 9c). In the Pearce (1983) multi-element plot, the samples display moderate negative anomalies in HFS elements such as Nb, P, Zr and Ti (Fig. 10c), which is characteristic of a subduction zone setting. All the lamprophyre dykes have high Zr/Y (4.0 to 9.0) and low Nb/Y (0.2 to 0.9), which is typical of within-plate tholeiitic basalt (Fig. 15a,b). In the Ti/1000 vs V diagram (Fig. 11), the lamprophyre dykes are characterized as MORB rocks.

On the Th/Yb and Ce/Yb vs Ta/Yb plots (Fig. 12a,b), the lamprophyre dykes plot as calc-alkaline to shoshonitic. On the La/Yb vs Th/Yb diagram (Fig. 13) they plot as continental

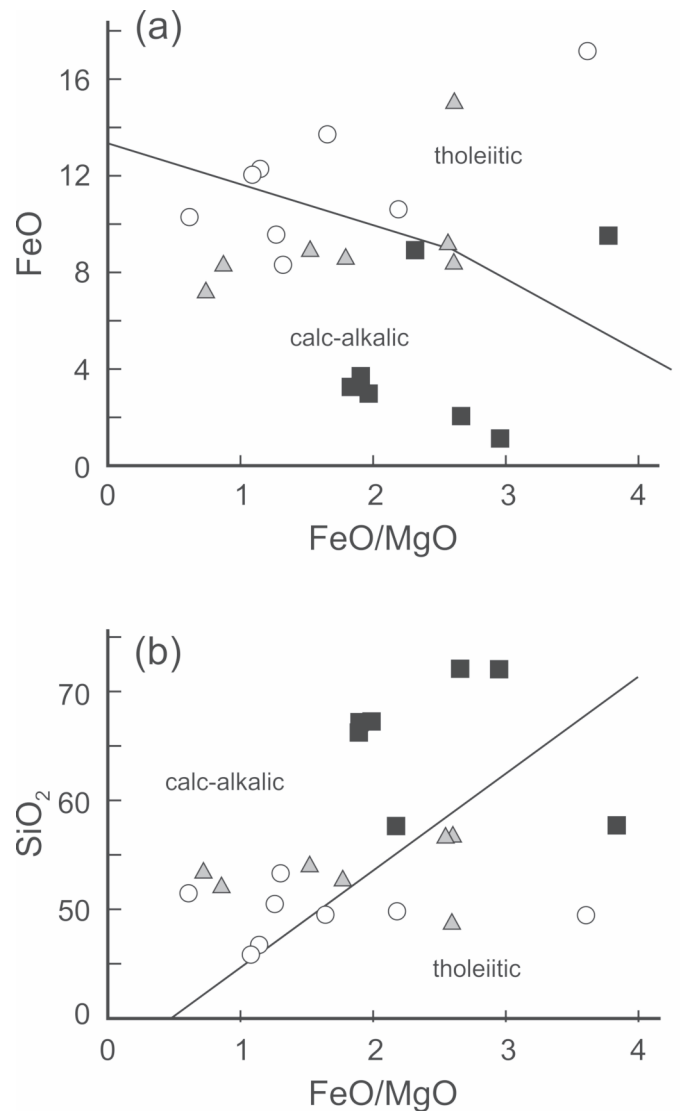


Figure 8. (a) FeO vs FeO_{tot}/MgO and (b) SiO₂ vs FeO_{tot}/MgO for the mafic and felsic rocks of the appinite suite in Donegal (dividing lines after Miyashiro 1974). Symbols as in Figure 7.

margin arc rocks, whereas on the Hf–Th–Ta discrimination diagram (Fig. 14), they plot as calc-alkaline basalt and have higher Th than typical island arc tholeiites (IAT), plotting close to the Th apex.

Sm–Nd Isotopic Analysis

Sm–Nd isotopic data for the Ar dara appinite suite and lamprophyre dykes are summarized in Table 2. Nd data are calculated at *t* = 435 Ma, based on the ⁴⁰Ar/³⁹Ar hornblende age determined for the suite. Sm and Nd values for rocks of the appinite suite and lamprophyre dykes lie in the ranges 1.99–6.86 ppm and 8.59–35.84 ppm, respectively. ¹⁴⁷Sm/¹⁴⁴Nd values range from 0.1071 to 0.1536 and ¹⁴³Nd/¹⁴⁴Nd values range from 0.512212 to 0.512490. Depleted mantle model ages (T_{DM}) were calculated according to DePaolo (1981, 1988). ε_{Nd}(*t*) values (*t* = 435 Ma) for the appinite suite range from –4.8 to +3.1 and T_{DM} values range from 805 Ma to 1710 Ma (Fig. 16).

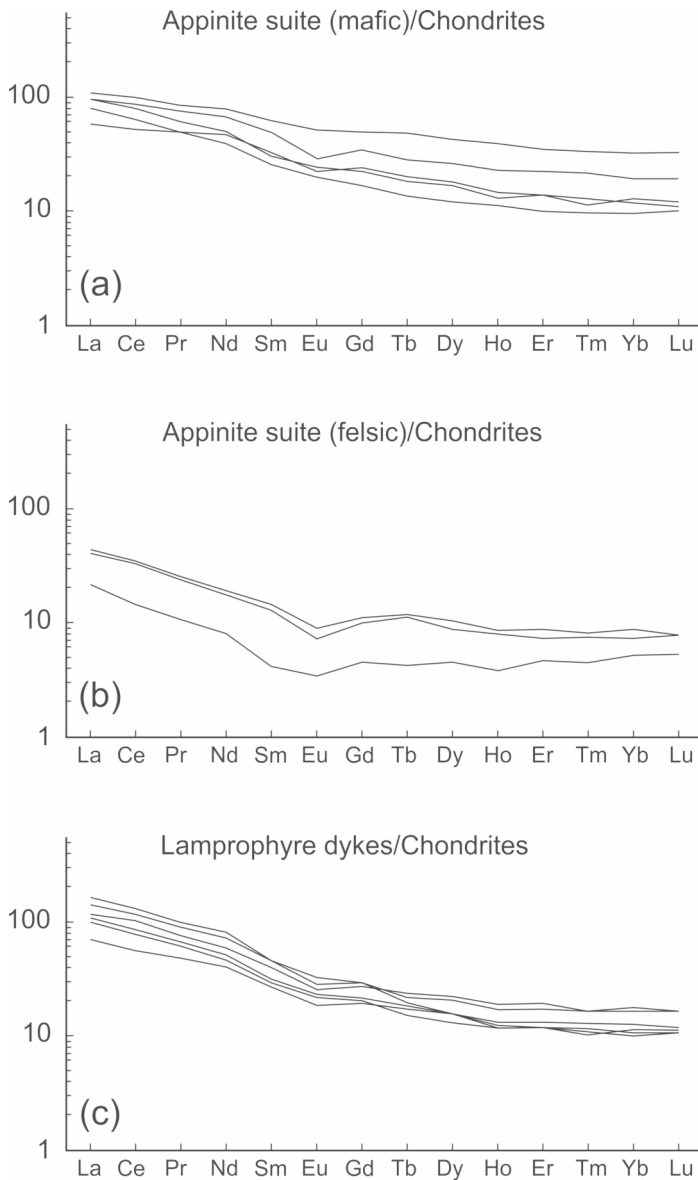


Figure 9. REE plots for representative samples of (a) mafic rocks of the appinite suite, (b) felsic rocks of the appinite suite and (c) the coeval lamprophyre dykes in Donegal (normalizing values from Sun and McDonough 1989).

DISCUSSION

The $^{40}\text{Ar}/^{39}\text{Ar}$ isotopic data closely constrain the timing of emplacement of the appinite suite. Given that the Ar closure temperature for hornblende is ca. 550°C (Harrison 1982) and that the host Dalradian metasedimentary rocks are at greenschist facies, the cooling ages obtained are interpreted to closely date igneous emplacement at ca. 434 Ma providing supporting evidence for the interpretation that the rocks of the appinite suite slightly predate the emplacement of the Ar dara granite, and significantly predate the intrusion of the Main Donegal Granite. This conclusion is similar to that interpreted for the association of granite with appinite and lamprophyre in the Scottish Caledonides (Rogers and Dunning 1991; Fowler and Henney 1996; Atherton and Ghani 2002; Neilson et al. 2009). In the latter case, slab break-off is thought to have induced

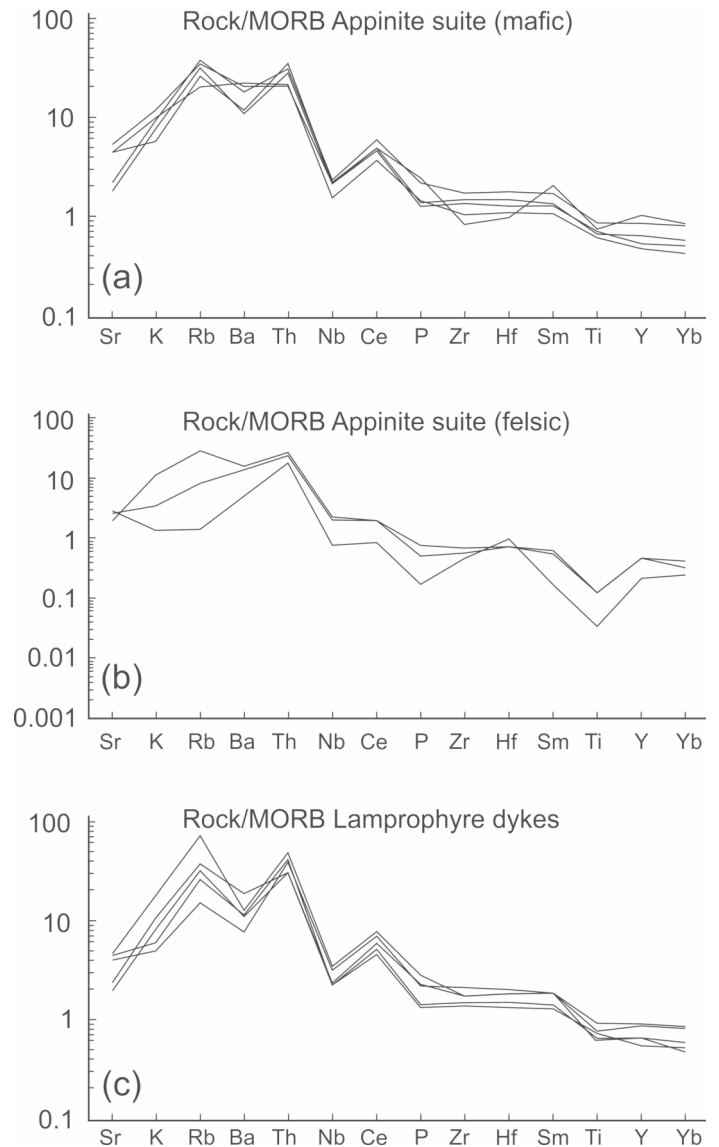


Figure 10. Multi-element plot for representative samples of (a) mafic rocks of the appinite suite, (b) felsic rocks of the appinite suite, and (c) the coeval lamprophyre dykes in Donegal (normalizing values from Pearce 1983).

asthenospheric upwelling, which induced melting of the lithospheric mantle producing an appinitic mafic layer that underplated the crust (Fig. 17). Heat from underplated mafic crust is inferred to have initiated the production of crustal melts and voluminous granitoid magmatism (e.g. Neilson et al. 2009). The granitoid bodies, in turn, may have been produced by fractional crystallization of appinitic parent magmas (von Blackenburg and Davies 1995; Fowler et al. 2008).

The appinite suite and lamprophyre dykes around the Ar dara pluton display similar chemical trends that, together with their field relationships, suggest they are comagmatic. For example, Al_2O_3 varies inversely and displays a systematic correlation with SiO_2 , and both rock groups show (i) negative correlations of SiO_2 with CaO, MgO, Fe_2O_3 , TiO_2 , V, Cr and Ni, and positive correlations of SiO_2 with K_2O and Na_2O , (ii) plot in the andesite to basaltic andesite fields on the Zr/Ti vs

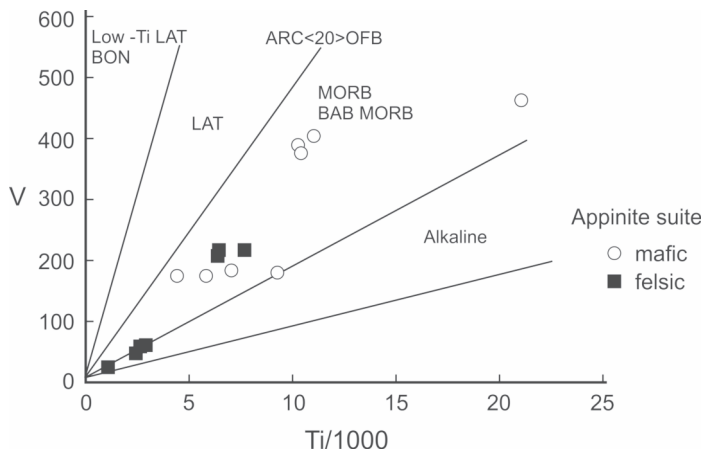


Figure 11. Ti/1000 vs V discrimination diagram for the appinitic suite rocks in Donegal (after Shervais 1982). IAT (island arc tholeiite); BON (boninitic basalt); BAB (back-arc basin basalt).

Nb/Y diagrams, and (iii) display very similar trends in REE plots, with elevated LREE and depleted HREE values. The same can be said for both groups of rocks on multi-element plots. Both groups additionally display negative anomalies in Nb, P, Zr and Ti, which are characteristics of a mantle source in subduction zone settings. This interpretation is supported by a 437 ± 5 Ma crystallization age of a lamprophyre dyke to the east of the Leannan Fault in Donegal (Kirkland et al. 2013).

The appinitic suite and lamprophyre dykes are also characterized by generally high Mg, Ni, Cr, Ba and V values (typical of mafic rocks), coupled with elevated LILE, and are moderately enriched in LREE with flat HREE profiles, indicating they were generated from partial melting of sub-continental lithospheric mantle. Both also show negative Nb and Ti anomalies, indicating either crustal contamination of the magmas or modification of the original mantle source by subduction zone fluids (e.g. Pearce 1996). If these anomalies were a result of crustal contamination, however, they should also be reflected in the Sm–Nd signature, resulting in a negative correlation between ϵ_{Nd} and parameters such as Ce/Yb and Th/Yb. As no such correlation is present (Fig. 18), the trace element contamination presumably occurred at the time of magma generation and so did not affect the isotopic system. Hence, the negative trace element anomalies probably reflect the composition of the metasomatized mantle source, implying that the low ϵ_{Nd} values of the appinitic suite and lamprophyre dykes directly reflect the composition of the mantle from which they were derived. But while the low ϵ_{Nd} values are consistent with derivation from the sub-continental lithospheric mantle, the range of ϵ_{Nd} values suggest the additional involvement of a juvenile source.

Rocks of the appinitic suite have been described from other orogenic belts of widely ranging age (e.g. Salmon 1998; Castro et al. 2002; Ye et al. 2008; Murphy 2013). Many are associated with major strike-slip faults (e.g. Murphy and Hynes 1990; Murphy 2013) including the appinitic suite of Donegal (e.g. Hutton 1982; Hutton and Alsop 1996; Kirkland et al. 2008). In this scenario, the Ar–Ar ages probably reflect the timing of

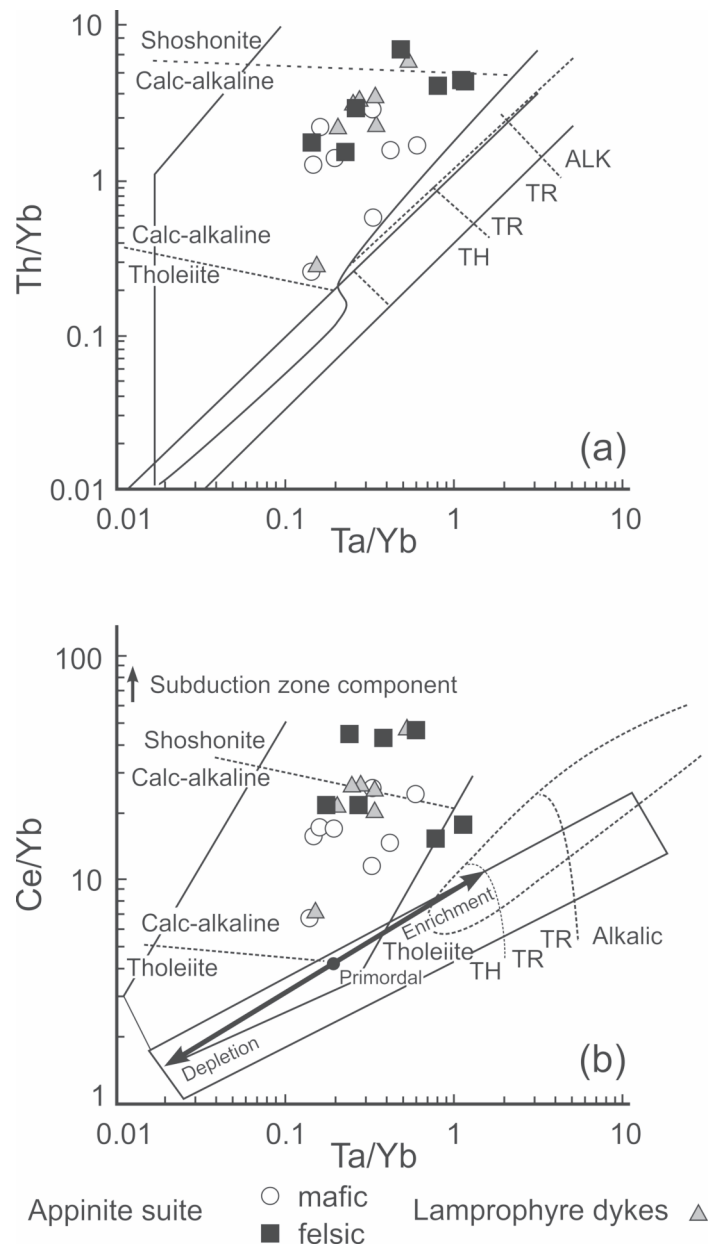


Figure 12. Log-log plots of (a) Th/Yb vs Ta/Yb and (b) Ce/Yb vs Ta/Yb (after Pearce 1982). TH (tholeiitic basalt); ALK (alkali basalt); TR (transition between tholeiitic and alkali basalt) for appinitic suite rocks in Donegal.

emplacement of appinitic magma into relatively shallow crustal levels along active strike-slip faults, rather than the duration of appinitic magmatism. Potential sources of the appinitic and lamprophyric magmas include melting of the sub-continental lithospheric mantle (SCLM), melting of individual metasomatic veins within the mantle (Mitchell 1995), and/or melting of metasomatized underplated mafic crust. Only limited fractionation occurred during magma ascent and the magma was chemically modified by (i) the addition of upper crustal material, or (ii) the addition of fluids derived from this upper crustal material.

Regional constraints indicate appinitic and lamprophyric magmatism was coeval with (or shortly followed) collision, for

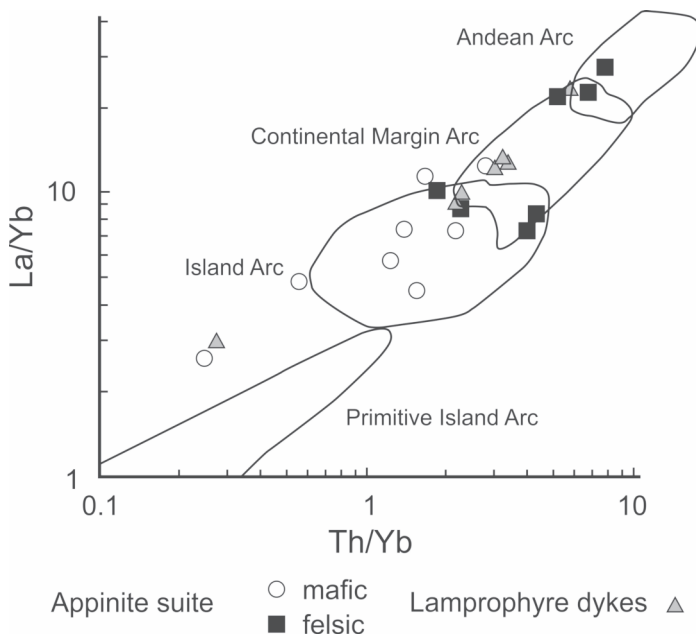


Figure 13. Log-log plot of La/Yb vs Th/Yb (after Condie 1989) for the appinite suite rocks of Donegal.

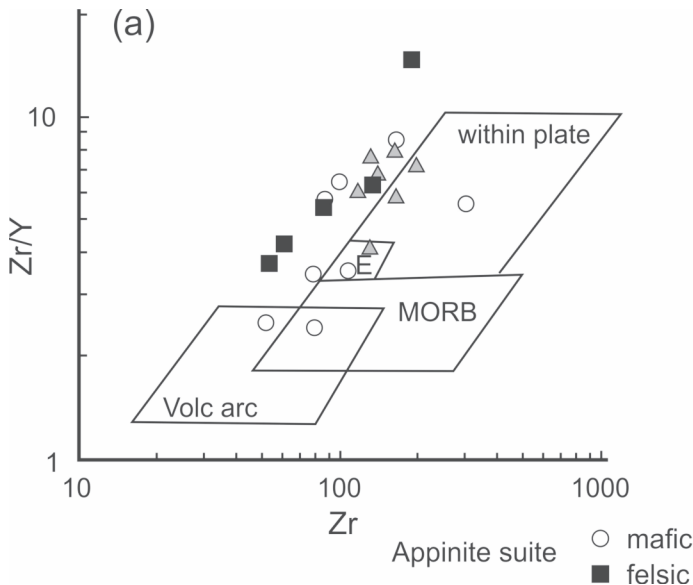


Figure 15. Zr vs Zr/Y discrimination diagram (after Pearce and Norry 1979) for the appinite suite rocks of Donegal. MORB (mid-ocean ridge basalt); E (enriched mid-ocean ridge basalt); VAB (volcanic arc basalt); WPB (within-plate basalt).

which slab break-off magmatism has been inferred (Ghani and Atherton 2008; Oliver et al. 2008; Neilson et al. 2009). Models proposed to account for appinitic and lamprophyric magmas in collisional settings ascribe the heat source to asthenospheric upwelling caused either by extensional tectonics associated with late orogenic events (Seyitoğlu and Scott 1996), such as lithospheric delamination, or by slab break-off (Dilek and Altunkaynak 2009). In both these scenarios, asthenospheric upwelling results in coeval melting of the overlying metasomatized SCLM and at least partial replacement of the SCLM (O'Reilly and Griffin 2006) by younger juvenile mantle (Downes 2001).

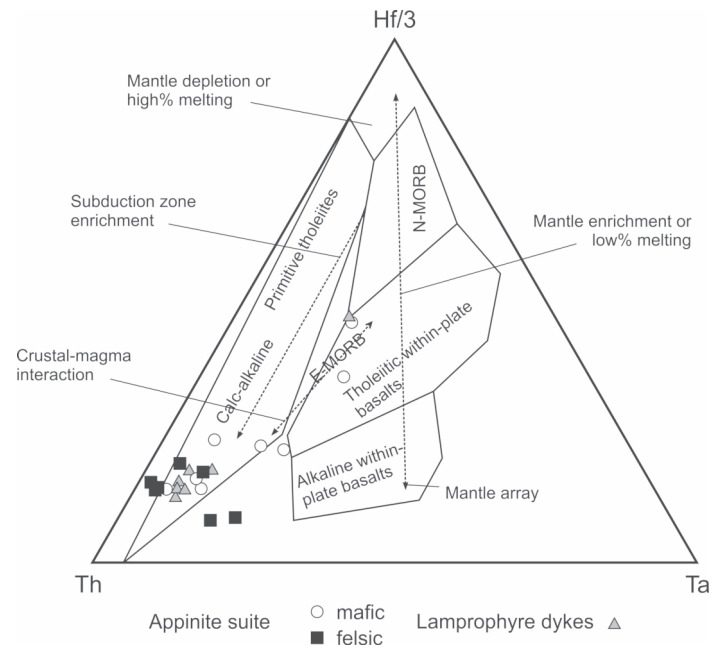
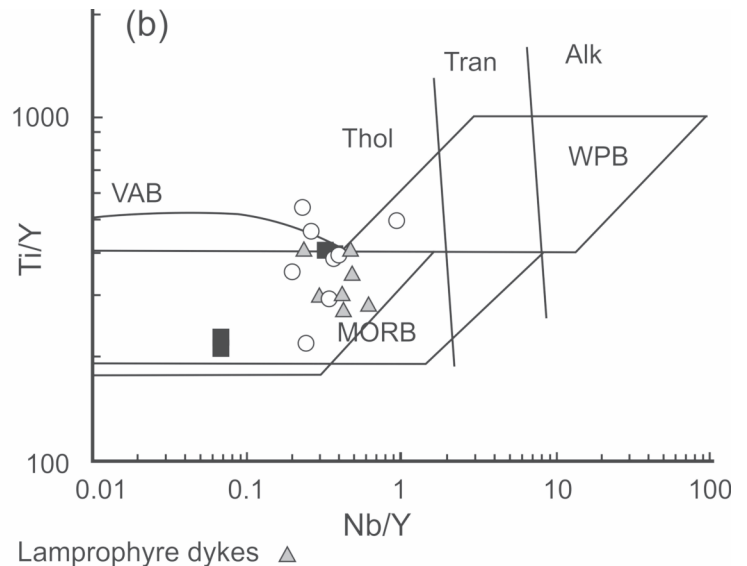


Figure 14. Triangular Hf/3-Ta-Th diagram (after Wood et al. 1979) for the appinite suite rocks of Donegal.



Regional Implications

The geochemical data presented here are consistent with previously published data for coeval post-collisional appinite plutons and lamprophyre dykes in the Scottish Highlands (Fowler 1988) whose emplacement also immediately preceded the emplacement of Caledonian granite plutons. These similarities suggest that processes responsible have regional significance. The origin of the rocks of the appinite suite and lamprophyre dykes in the Caledonides has been related specifically to the closure of the Iapetus Ocean, which was immediately followed by subduction of Avalonia beneath Laurentia (Bowes and Košler 1993). Slab break-off is likely to have been a conse-

Table 2. Sm–Nd isotope, ϵ_{Nd} and T_{DM} age data for the mafic and intermediate–felsic rocks of the appinite suite and coeval lamprophyre dykes peripheral to the Ardara pluton, Donegal, Ireland.

Sample	Nd (ppm)	Sm (ppm)	$^{147}\text{Sm}/^{144}\text{Nd}$	$^{143}\text{Nd}/^{144}\text{Nd}$	2s	ϵ_{Nd} (0 Ma)	ϵ_{Nd} (435 Ma)	T_{DM}
DWH-003	21.71	4.38	0.1219	0.512347	6	-5.7	-1.5	1156
DMH-003	35.84	6.86	0.1158	0.512454	7	-3.6	0.9	922
FEL-001	18.41	4.07	0.1335	0.512478	7	-3.1	0.4	1076
LMB-009	31.75	6.43	0.1225	0.512318	7	-6.2	-2.1	1210
LBB-002	24.05	4.88	0.1226	0.512257	7	-7.4	-3.3	1311
INK-004	20.25	5.14	0.1536	0.512365	5	-5.3	-2.9	1710
AGR-001	17.90	3.67	0.1239	0.512381	6	-5.0	-1.0	1124
GLE-001	14.76	3.65	0.1496	0.512336	7	-5.9	-3.3	1673
GLE-004	21.18	5.25	0.1500	0.512339	7	-5.8	-3.2	1676
LBB-003	22.03	4.71	0.1294	0.512603	6	-0.7	3.1	805
INK-008	8.59	1.99	0.1400	0.512231	8	-7.9	-5.8	1673
FEL-005	20.39	3.61	0.1071	0.512212	7	-8.3	-3.3	1189

($^{143}\text{Nd}/^{144}\text{Nd}$) is adjusted from the deviation to JNdi-1 Standard (accepted value = 0.512115), mean measured value of the standard gives: $^{143}\text{Nd}/^{144}\text{Nd}$ (0.512105) Std (0.000008) n (22)

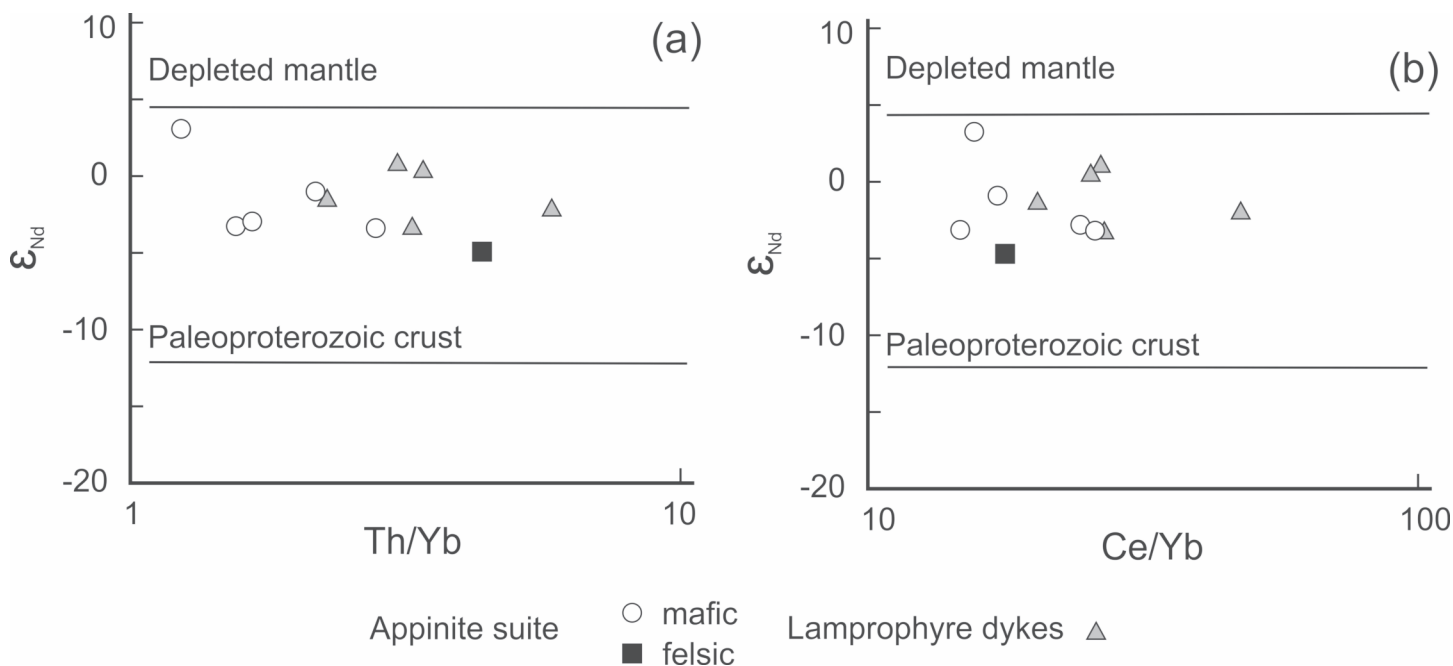


Figure 16. Nd isotope evolution diagram for the appinite suite rocks of Donegal. Values shown on the depleted mantle curve are T_{DM} ages (DePaolo 1981, 1988).

quence of this event. The magmatic association related to this collision (high-K, calc-alkaline, lamprophyric compositions) and the accompanying granitoid magmatism are best interpreted as direct responses to slab-breakoff, an association recently suggested for some late-stage Cordilleran batholiths (e.g. Hildebrand and Whalen 2014a, b). Within the Caledonian orogen in Britain, the association in time and space of appinitic, lamprophyric and voluminous granitic magmas is one in which the mantle components are chemically and isotopically similar and the magmas of all three rock groups are genetically related (Rock and Hunter 1987; Fowler and Henney 1996).

The felsic components include large batholiths that were probably derived by fractionation of shoshonitic magma

(Fowler 1988). Other appinitic suites include some, but not all, of these features (Murphy 2013). Some common tectonic traits of appinitic rocks include a tendency to be emplaced soon after the cessation of subduction, and the important role of deep crustal faults as conduits for magmas of various compositions to rise towards the surface. These conduits provide the setting for magmas of diverse composition to mix and mingle (Murphy 2013). The magmas are quite similar to those of lamprophyre dykes and, together with ultrapotassic magma, are considered to have a parental source in potassium-rich phlogopite-bearing mantle. The appinitic and lamprophyric rocks of Donegal may therefore be an example of the products of collisional asthenospheric upwelling that occurred in this area

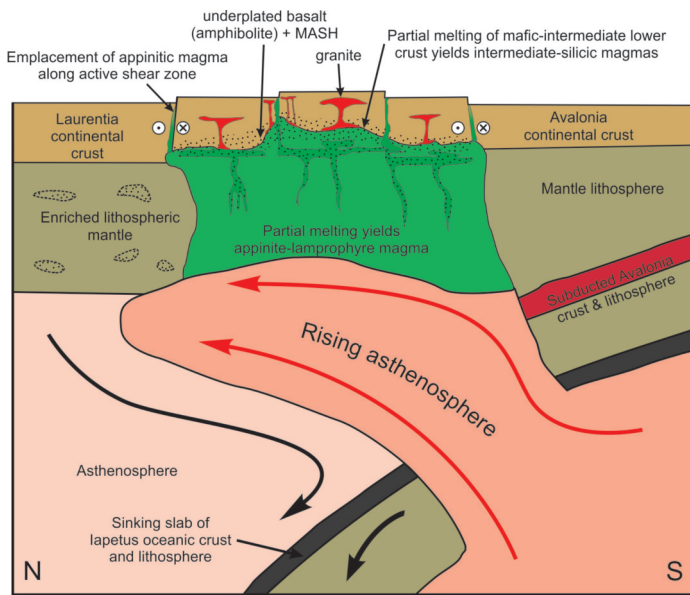


Figure 17. Schematic illustration (modified after Neilson et al. 2009) of a model for magmatism in the Caledonides of Scotland and NW Ireland relating the origin of appinite magma through slab break-off and the consequent upwelling of hot asthenosphere. This upwelling is thought to have caused partial melting of the lithospheric mantle, producing an appinitic-lamprophyre mafic layer that underplated the crust and initiated the production of voluminous granitoid magma. Magma ascent was likely facilitated by major faults (e.g. Jacques and Reavy 1994).

during the Caledonian orogeny. Whether these occurrences can be attributed to an orogen-wide delamination event or to localized zones where slab break-off occurred would require further investigation.

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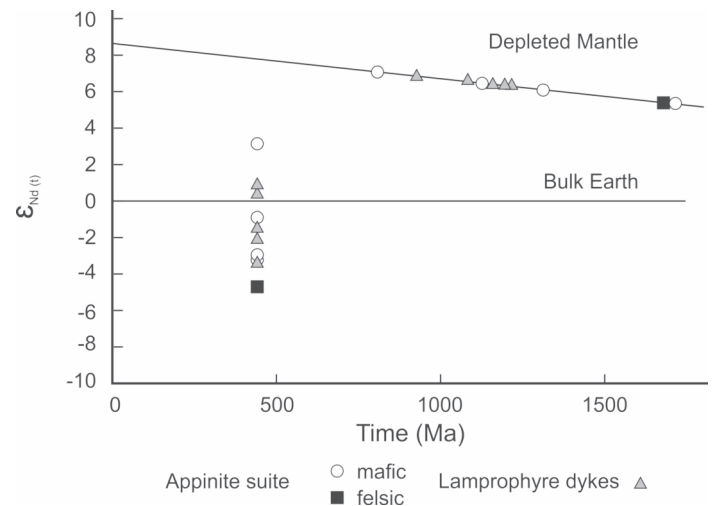


Figure 18. Plots of (a) ϵ_{Nd} vs Th/Yb, and (b) ϵ_{Nd} vs Ce/Yb (hypothetical evolution of Paleoproterozoic crust shown as reference).

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